

Assessment of Radiogenic Heat Production Due To ^{238}U , ^{232}Th and ^{40}K In Surface Soils of Ortum Region, West Pokot County, Kenya

¹Mukanda Kere Wanyama, ²Michael Nakitare Waswa, ³Felix Omonya Wanjala

¹ Department of Science, Technology & Engineering; Kibabii University, 1699-50200 Bungoma

²Department of Science, Technology & Engineering; Kibabii University, 1699-50200 Bungoma

³IAEA, Vienna, Austria

ABSTRACT

The activity concentrations of ^{238}U , ^{232}Th and ^{40}K in Bq/kg were determined using a HPGe detector at the Institute of Nuclear Science and Technology at the University of Nairobi (Kenya). The activity concentrations were then converted into ppm (^{238}U and ^{232}Th) and % for ^{40}K that were used for computing Radiogenic Heat Production (RHP) for individual radionuclides and the overall RHP contributed by the three radionuclides. The average activity concentrations were 3.198753ppm, 13.42265ppm and 1.372154% for ^{238}U , ^{232}Th and ^{40}K respectively. The average RHP for the individual radionuclides was found to be $0.239883\mu\text{W}/\text{m}^3$, $0.2658\mu\text{W}/\text{m}^3$ and $0.036378\mu\text{W}/\text{m}^3$ for ^{238}U , ^{232}Th and ^{40}K respectively. The overall RHP contribution by the three radionuclides was found as $0.540764\mu\text{W}/\text{m}^3$. All the RHP values for the individual radionuclides including their overall contribution was below the world RHP value of $1\mu\text{W}/\text{m}^3$. The findings of this study will provide benchmark information on RHP status of the study area that can be extended to other areas to enable possible geothermal energy exploitation.

KEYWORDS

radiogenic heat production;
natural radioactivity; elemental
concentration

INTRODUCTION

Radiogenic heat production refers to the thermal energy resulting from radioactive decay. The radioactive disintegration is due to gravitational contraction and decay of unstable radioisotopes [1]. RHP is associated with the radioisotopes: ^{238}U , ^{232}Th and singly occurring ^{40}K due to their enrichment relative to other radionuclides [2]. Besides the heat of the earth, the RHP due to ^{238}U , ^{232}Th and ^{40}K constitutes the biggest internal source of heat. According to [3], the annual radiogenic heat production in the earth is 6.3×10^{20} J which is greatly more than twice the world primary energy. Most mass of the radionuclide is converted into heat during radioactive decay [4]. The radionuclides of ^{238}U , ^{232}Th and ^{40}K are associated with long half-lives: ^{40}K (half-life = 1.277×10^9 years), ^{238}U (half-life = 4.468×10^9 years) and ^{232}Th (half-life = 1.405×10^{10} years). Most of the significant emissions from the radionuclides are converted into heat especially within the rocks [4,5]. The energy emitted by all the decay consisting of kinetic energy of the emitted particles and gamma radiation is absorbed and eventually transformed into heat [7]. The large energy source from the earth alone is greater than the global requirement predicted by [6] for the year 2030. Thus, if this energy is tapped to a large extent then it may take care of the demands of the primary energy for the rest of the years. According to the study by [8], RHP is an important parameter in geothermal exploitation studies. It is also essential in explaining continental thermal flux density for providing information on the temperature and structure of the lithosphere [9]. Several studies have been done globally on radiogenic heat production due to natural radioactivity. According to a study by [10] on RHP and excess life time cancer risk of the sand in India some samples registered higher RHP values compared with the world average value of $1\mu\text{W}/\text{m}^3$ although the average for all the samples was below $1\mu\text{W}/\text{m}^3$.

From the study done by [11] on migmatites, biotite gneisses using gamma ray spectroscopy, some of the rocks revealed high levels of natural radioactivity and hence some of the RHP values for the samples were high. In addition, [12] did a study on RHP in crustal quarry rocks where the mean RHP ranged from 0.549 to 3.122 $\mu\text{W}/\text{m}^3$ with Uranium contributing the highest RHP with a correlation coefficient of 83.09%. In the East Africa, there are scanty researches on RHP. In Uganda, [13] did study on RHP in river sediments from Gulu and Amum districts. The RHP values ranged from 2.5 \pm 0.6 $\mu\text{W}/\text{m}^3$ to 10.2 \pm 2.6 $\mu\text{W}/\text{m}^3$ $\mu\text{W}/\text{m}^3$ with an average of 4.5 $\mu\text{W}/\text{m}^3$. The RHP value was higher than the usual average value of 1 $\mu\text{W}/\text{m}^3$.

MATERIALS AND METHODS

Study area& Sample Preparation

The study was done in Ortum region, Kapenguria constituency, West Pokot County. The area studied was 5km², the study area lies between longitude 35°15'E and 35°20'E and latitude 1°10'N and 1°30'N and an altitude of about 1550m above sea level as shown in Figure 1. Ortum region is surrounded by Chereng'anyi and Ortum hills with high presence of granitic and silicon rocks that are associated with high levels of background radiations [14]. Most of the population in Ortum depends on Agriculture, quarrying, gold mining and building and construction materials.

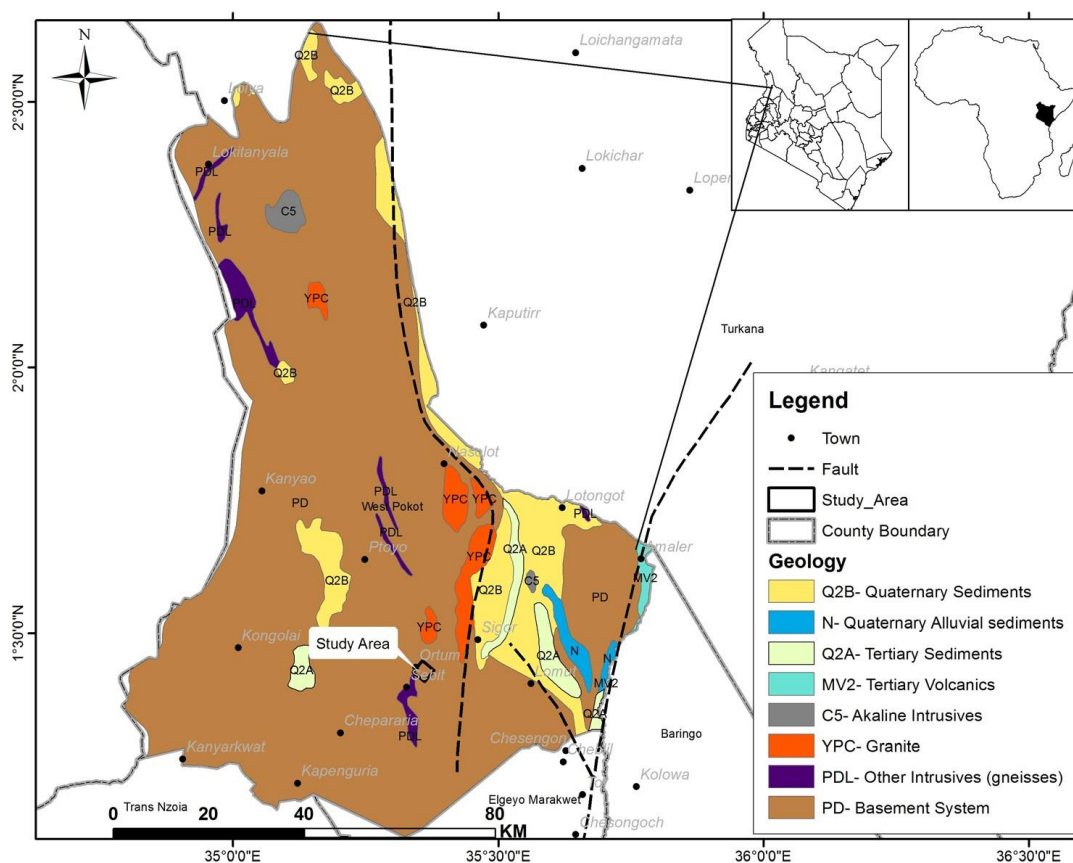


FIGURE 1: Map showing the study area.

A total of 35 soil samples each of about 1kg from the surface at different points (Figure 2) were collected for assessment of ²³⁸U, ²³²Th and ⁴⁰K levels and their radioelement concentrations. The samples were enclosed in containers properly labeled to avoid mix up. The soil sample were collected using a soil auger, trowel and a panga. The foreign materials: pebbles, roots and grass were removed. The samples were then dried in the oven for 24 hours at 110 OC and then crushed into fine powder and sieved through 100 μm sieve. About 350g of each of the soil samples were put and sealed in air tight plastic containers and kept for a minimum of 30 days to allow for secular equilibrium between ²²⁶Ra, ²³²Th and their progenies [15]. The HPGc detector was then used to determine the activity concentrations of the radionuclides of individual samples. Figure 2 below shows the sampling points where the samples were obtained.

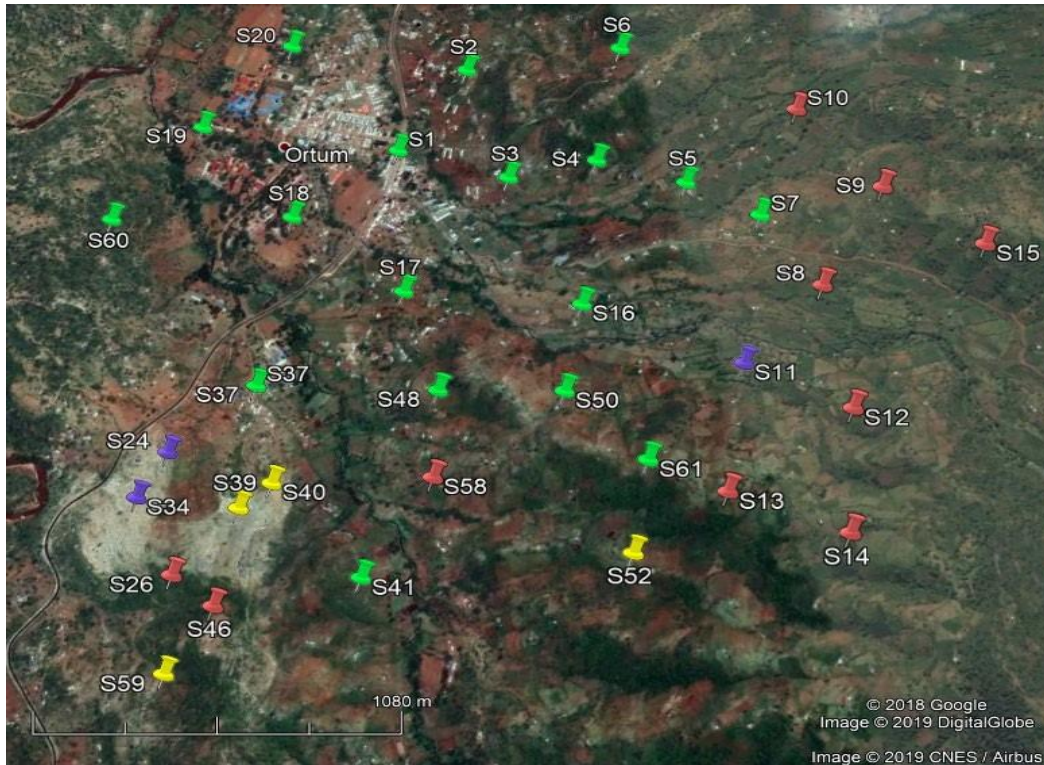


FIGURE 2: Surface soil sampling points.

Experimental Techniques

The HPGe detector at the institute of Nuclear Science and Technology (University of Nairobi) was used to determine the activity concentrations of ²³⁸U, ²³²Th and ⁴⁰K in the soil and rock samples. The efficiency calibration was done using standard reference materials: RGU - 1, RGTh - 1 and RGK - 1 provided by the IAEA with known activity concentrations using equation 1[16].

$$\epsilon = \frac{N}{A \times P \times M \times T} \tag{1}$$

where ϵ is the efficiency, A is the activity concentration, N is the net area under the photo peak, T is the counting time, P is the emission probability and m is the mass of the sample.

The energy calibration was done by relating the channel number with the photo energy using gamma lines of ²¹⁴Pb, ²¹⁴Bi, ²²⁸Ac and ⁴⁰K with energies 351Kev, 609Kev, 911Kev and 1460Kev. The energy was evaluated using equation 2 [16].

$$E = a (ch) + b \tag{2}$$

where E is the energy, a and b are constants and ch is the channel number.

Determination of Activity Concentrations

The activity concentrations of the individual radionuclides were determined using equation 3 [16].

$$\frac{M_s A_s}{I_s} = \frac{M_r A_r}{I_r} \tag{3}$$

where M_s is the mass of the sample, A_s is the activity concentration of the sample, I_s is the intensity of the sample, M_r is the mass of the standard sample, A_r is the activity concentration of standard reference material and I_r is the intensity of the reference material.

Radiogenic Heat Production (RHP)

The Radiogenic Heat Production was computed by converting the activity concentrations of the individual samples into ppm using equations 4 [17]

$$\left. \begin{aligned} 1\text{ppm of } 238_U &\equiv 12.35 \frac{\text{Bq}}{\text{Kg}} \\ 1\text{ppm of } 232_{Th} &\equiv 4.06 \frac{\text{Bq}}{\text{Kg}} \\ 1\% \text{ of } 40_K &\equiv 313 \frac{\text{Bq}}{\text{Kg}} \end{aligned} \right\} \quad (4)$$

The sum total of the individual contributions of C_U , C_{Th} and C_K in the soil and rock samples gives the overall heat production given by equation 5 [18].

$$RHP \left(\frac{\mu\text{W}}{\text{m}^3} \right) = 10^{-5} \rho [9.52C_U(\text{ppm}) + 2.56C_{Th}(\text{ppm}) + 3.48C_K(\%)] \quad (5)$$

where ρ is the density of the individual samples.

RESULTS AND DISCUSSIONS

Activity concentrations of the radionuclides and associated RHP in the surface soil samples

The activity concentrations of ^{238}U , ^{232}Th and ^{40}K in the 35 soil samples were determined using the HPGe detector following equation 3. Thereafter the activity concentrations were converted into individual radioelement concentrations in ppm using the relations 4 for use in computing the RHP using equation 5. The table 1 below shows the activity concentrations of the three radionuclides: ^{238}U , ^{232}Th and ^{40}K in ppm (for ^{238}U and ^{232}Th) and % for ^{40}K and the associated radiogenic heat produced due to the three radionuclides.

TABLE 1: Activity concentration of the samples and associated RHP.

Samples	U-238 (ppm)	Th-232 (ppm)	K-40 (%)	U-238 RHP($\mu\text{W}/\text{m}^3$)	Th-232 RHP($\mu\text{W}/\text{m}^3$)	K-40 RHP($\mu\text{W}/\text{m}^3$)	Total RHP($\mu\text{W}/\text{m}^3$)
S1	2.069636	11.46305	0.543035	0.153249	0.228249	0.014699	0.396197
S2	3.523887	13.76601	1.08901	0.260932	0.274104	0.029477	0.564513
S3	2.808907	11.23892	0.985719	0.20799	0.223786	0.026681	0.458456
S4	2.071255	9.263547	0.473674	0.153369	0.184453	0.012821	0.350643
S5	1.653441	8.273399	0.866358	0.122416	0.164737	0.02345	0.310603
S6	2.11498	9.288177	0.807284	0.156607	0.184943	0.021851	0.363401
S7	2.006478	8.800493	0.788243	0.148573	0.175233	0.021336	0.345141
S8	2.877733	12.78818	1.0123	0.213086	0.254634	0.0274	0.495121
S9	4.263158	17.54433	3.257412	0.315672	0.349337	0.08817	0.753179
S10	4.451012	17.68227	3.200607	0.329582	0.352084	0.086632	0.768298
S11	3.145749	15.45074	1.820511	0.232932	0.30765	0.049277	0.589859
S12	2.553036	11.42365	0.995272	0.189044	0.227464	0.026939	0.443447
S13	2.823482	11.97537	0.884824	0.209069	0.23845	0.02395	0.471469
S14	3.215385	13.00739	1.311406	0.238088	0.258999	0.035496	0.532584
S15	2.039676	8.997537	0.730639	0.151031	0.179156	0.019777	0.349964
S16	5.417814	17.94828	2.795719	0.40117	0.35738	0.075673	0.834224
S17	2.235628	9.837438	0.920032	0.165541	0.19588	0.024903	0.386323
S18	3.866397	15.37438	1.448658	0.286293	0.30613	0.039211	0.631635
S19	2.259109	8.980296	0.904377	0.167279	0.178813	0.024479	0.370571
S20	2.484211	11.04926	0.922204	0.183947	0.220009	0.024962	0.428918
S24	3.783806	15.85961	1.588466	0.280178	0.315791	0.042996	0.638965
S26	4.77247	18.30542	2.138403	0.353385	0.364492	0.057881	0.775758

Samples	U-238 (ppm)	Th-232 (ppm)	K-40 (%)	U-238 RHP($\mu\text{W}/\text{m}^3$)	Th-232 RHP($\mu\text{W}/\text{m}^3$)	K-40 RHP($\mu\text{W}/\text{m}^3$)	Total RHP($\mu\text{W}/\text{m}^3$)
S34	2.394332	10.31034	0.960863	0.177292	0.205296	0.026008	0.408596
S37	2.65668	11.94335	0.919553	0.196718	0.237812	0.02489	0.45942
S39	2.487449	10.53695	0.847572	0.184187	0.209808	0.022942	0.416937
S40	4.287449	17.88424	1.324952	0.317471	0.356105	0.035863	0.709439
S41	2.660729	11.81773	0.981981	0.197018	0.235311	0.02658	0.458908
S46	2.655061	11.00246	0.725911	0.196598	0.219078	0.019649	0.435324
S48	5.018623	20.50493	2.325815	0.371612	0.408288	0.062954	0.842853
S50	4.100405	17.94089	1.759553	0.303621	0.357233	0.047627	0.708481
S52	4.321457	17.61823	2.245559	0.319989	0.350808	0.060782	0.731579
S58	2.559514	10.78571	0.913514	0.189523	0.214762	0.024726	0.429011
S59	2.909312	12.60099	1.260256	0.215425	0.250907	0.034112	0.500443
S60	3.22915	15.11576	1.107732	0.239107	0.30098	0.029983	0.570071
S61	5.491498	20.99261	2.181214	0.406627	0.417998	0.05904	0.883665
Min	1.653441	8.273399	0.473674	0.122416	0.164737	0.012821	0.310603
Max	5.491448	20.99261	3.257412	0.406627	0.417998	0.08817	0.883665
Av	3.198753	13.42265	1.372154	0.239883	0.2658	0.036378	0.540764

The results of the activity concentrations of the three radionuclides: ^{238}U , ^{232}Th and ^{40}K showed a spatial distribution pattern. The radiation levels of ^{238}U , ^{232}Th and ^{40}K ranged from 1.653441ppm to 5.491448 ppm, 8.273399 ppm to 20.99261ppm and 0.473674% to 0.406627% respectively. The average concentrations of the three radionuclides were 3.198753 ppm, 13.42265 ppm and 1.372154% for ^{238}U , ^{232}Th and ^{40}K respectively that were higher than the world average (18). ^{232}Th had the highest deviation from world average. The higher activity concentrations could be attributed to the anthropogenic activities especially agricultural like the use of phosphatic fertilizers for increased yields. The results from this study are similar to those of [19], [20] and [21]. The activity concentrations of ^{238}U and ^{232}Th were also higher than those from the study by (23) The individual RHP values for ^{238}U , ^{232}Th and ^{40}K ranged from 0.122416 $\mu\text{W}/\text{m}^3$ to 0.406627 $\mu\text{W}/\text{m}^3$, 0.164737 to 0.417998 $\mu\text{W}/\text{m}^3$, 0.012821 $\mu\text{W}/\text{m}^3$ to 0.08817 $\mu\text{W}/\text{m}^3$ respectively. The average radiogenic heat production was 0.239883 $\mu\text{W}/\text{m}^3$, 0.2658 $\mu\text{W}/\text{m}^3$ and 0.036378 $\mu\text{W}/\text{m}^3$ for ^{238}U , ^{232}Th and ^{40}K respectively. The overall average RHP contributed by all the three radionuclides was 0.540764 $\mu\text{W}/\text{m}^3$. The results from this study agrees with that of [7] where the study obtained an average $0.48 \pm 0.1 \mu\text{W}/\text{m}^3$. The results, however, deviates from those of [13] where they obtained an average RHP of $4.5 \pm 1.1 \mu\text{W}/\text{m}^3$. The difference in the results can be attributed to the different geological set ups of the two study areas considering the type of rocks from which the soils originated. For complete discussions of the results, correlation graphs were drawn to establish some relationships as shown in Figures 3a, 3b and 3c. In Figure 3a of ^{232}Th and ^{238}U the correlation coefficient is $R^2 = 0.934$ showing a very strong correlation. The correlation coefficients of 3b and 3c are $R^2 = 0.7279$ and $R^2 = 0.6912$ respectively both lying between 0.60 and 0.79 showing a strong correlation. The individual radiogenic heat produced for the three radionuclides were further represented in form of a graph as shown in Figure 4. From the figure 4, it can be seen that ^{232}Th and ^{238}U are the major contributors to the RHP. The contribution by ^{40}K is very small.

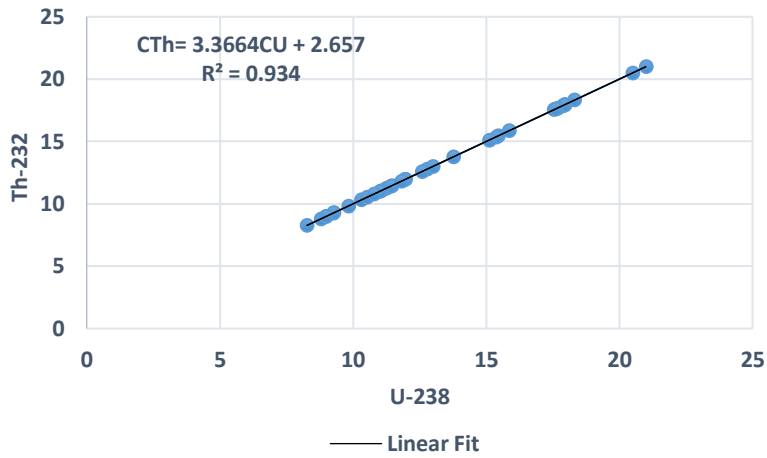


FIGURE 3a: Correlation between ^{232}Th and ^{238}U .

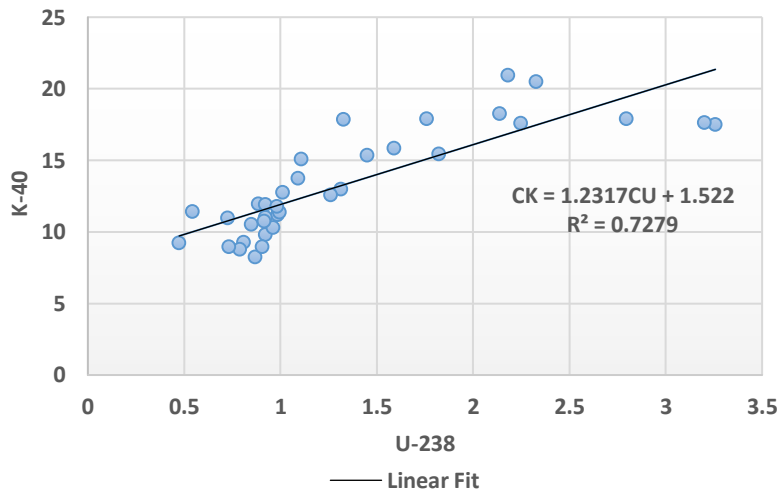


FIGURE 3b: Correlation between ^{40}K and ^{238}U .

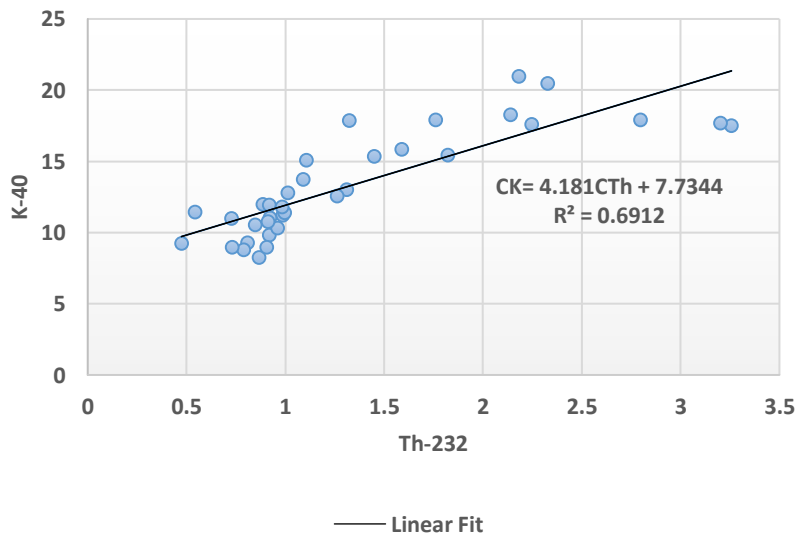


FIGURE 3c: Correlation between ^{40}K and ^{232}Th .

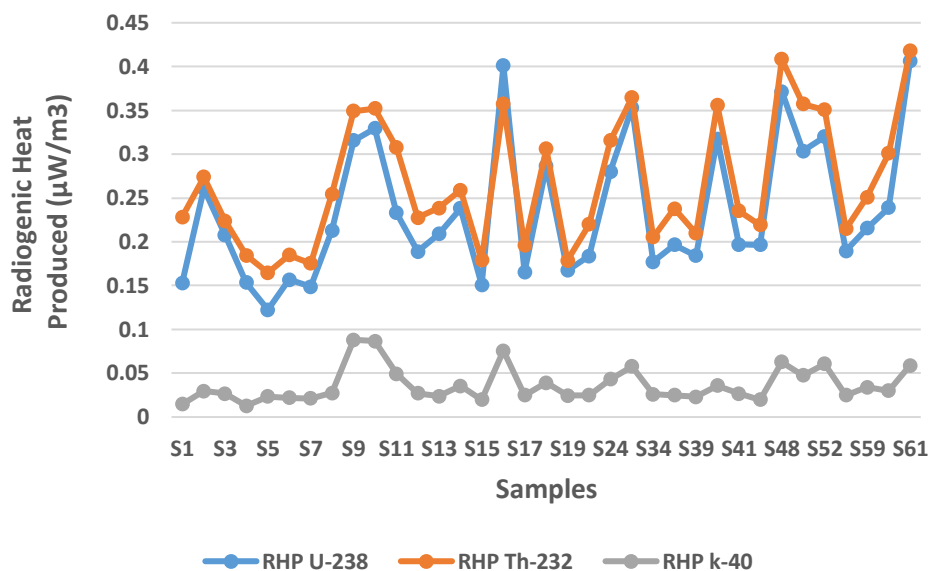


FIGURE 4: RHP comparison for the radionuclides.

CONCLUSIONS

An assessment on RHP due to ^{238}U , ^{232}Th and ^{40}K in surface soils of Ortum region has been done in this study. A HPGe detector was used to determine the activity concentrations of the three radionuclides in Bq/kg that were then converted into ppm (^{238}U and ^{232}Th) and % for ^{40}K . The average activity concentrations were 3.198753ppm, 13.42265ppm and 1.372154% for ^{238}U , ^{232}Th and ^{40}K respectively. the average RHP for the individual radionuclides was found as 0.239883 $\mu\text{W}/\text{m}^3$, 0.2658 $\mu\text{W}/\text{m}^3$ and 0.036378 $\mu\text{W}/\text{m}^3$ for ^{238}U , ^{232}Th and ^{40}K respectively. the overall RHP contribution by the three radionuclides was found as 0.540764 $\mu\text{W}/\text{m}^3$. All the RHP values for the individual radionuclides including their overall contribution was below the world value of 1 $\mu\text{W}/\text{m}^3$. This study being the first of the kind at the study area, it will provide benchmark information that can be extended to other areas. Further studies can be done on soils at different depths and on rocks to have comprehensive data base information on RHP values that may prompt possible geothermal energy exploitation.

ACKNOWLEDGEMENTS

The authors thank the staff at the Institute of Nuclear Science and Technology for assisting in the running of the samples. The authors are also grateful to the people of Ortum for their cooperation during the collection of the samples.

CONFLICT OF INTEREST

The authors declare no conflict of interest

REFERENCES

- [1] Clauser, C. (2009). Heat transport processes in the Earth's crust. *Surveys in Geophysics*, 30, 163-191.
- [2] Mohamed, H., Mizunaga, H., Abou Ashour, N. M., Elterb, R. A., Elalfy, I. M., & Elsayed, A. S. (2017). Radiogenic heat production in Rudeis Formation, Lower Miocene, Belayim marine oil field, Gulf of Suez, Egypt. *Exploration Geophysics*, 48(4), 512-522.
- [3] Jaupart, C., Labrosse, S., Lucazeau, F., & Mareschal, J. C. (2007). 7.06-temperatures, heat and energy in the mantle of the earth. *Treatise on geophysics*, 7, 223-270.
- [4] Bückler, C., Jarrard, R. D., & Wonik, T. (2001). Downhole temperature, radiogenic heat production, and heat flow from the CRP-3 drillhole, Victoria Land Basin, Antarctica. *Terra Antarctica*, 8(3), 151-160.
- [5] Ali, S., & Orazulike, D. M. (2010). Well logs-derived radiogenic heat production in the sediments of the Chad Basin, NE Nigeria. *Journal of Applied Sciences*, 10(10), 786-800.

- [6] Nichols, A. L., Verpelli, M., & Aldama, D. L. (2008). *Handbook of nuclear data for safeguards: database extensions, August 2008* (Vol. 818). IAEA.
- [7] Okeyode, I. C. (2012). Radiogenic heat production due to natural radionuclides in the sediments of Ogun River, Nigeria. *Geography*, 2(10).
- [8] Abbady, A. G., & Al-Ghamdi, A. H. (2018). Heat production rate from radioactive elements of granite rocks in north and southeastern Arabian Shield Kingdom of Saudi Arabia. *Journal of Radiation Research and Applied Sciences*, 11(4), 281-290.
- [9] Rybach, L. (1988). Determination of heat production rate. *Handbook of terrestrial heat-flow density determination*, 125-142.
- [10] Murugesan, S., & Ravichandran, S. (2023). Radioactive heat production rate and excess lifetime cancer risk of sand from two major rivers in India—A comparative study. *International Journal of Radiation Research*, 21(1), 117-124.
- [11] Akingboye, A. S., Ogunyele, A. C., Jimoh, A. T., Adaramoye, O. B., Adeola, A. O., & Ajayi, T. (2021). Radioactivity, radiogenic heat production and environmental radiation risk of the Basement Complex rocks of Akungba-Akoko, southwestern Nigeria: insights from in situ gamma-ray spectrometry. *Environmental Earth Sciences*, 80, 1-24.
- [12] Ogunsanwo, F. O., Adepitan, J. O., Ayanda, J. D., Giwa, K. W., Falayi, E. O., & Adejimi, A. I. (2021). Radiogenic heat production in crustal quarry rocks of Ogun State, south-western, Nigeria. *Environmental Earth Sciences*, 80, 1-15.
- [13] Olanyaa, A., Okello, D., Oruruc, B., & Kisolod, A. Natural Radioactivity Levels and Radiogenic Heat Production in River Sediments from Gulu and Amuru Districts, Northern Uganda.
- [14] Ferreira, A. D. O., & Pecequilo, B. R. (2019). Dose rates evaluation of some granitic rocks from the Paraná state. In *Proceedings of the INAC 2019: international nuclear atlantic conference. Nuclear new horizons: fueling our future*.
- [15] Nwankwo, C. U., Ogundare, F. O., & Folley, D. E. (2015). Radioactivity concentration variation with depth and assessment of workers' doses in selected mining sites. *Journal of Radiation Research and Applied Sciences*, 8(2), 216-220.
- [16] Otwoma, D., Patel, J. P., Bartilol, S., & Mustapha, A. O. (2013). Estimation of annual effective dose and radiation hazards due to natural radionuclides in mount Homa, southwestern Kenya. *Radiation protection dosimetry*, 155(4), 497-504.
- [17] Čermák, V., Bodri, L., & Rybach, L. (1991). Radioactive heat production in the continental crust and its depth dependence. *Terrestrial heat flow and the lithosphere structure*, 23-69.
- [18] United Nations Scientific Committee on the Effects of Atomic Radiation. (2000). *Sources and Effects of Ionizing Radiation, United Nations Scientific Committee on the Effects of Atomic Radiation (UNSCEAR) 2000 Report, Volume I: Report to the General Assembly, with Scientific Annexes-Sources*. United Nations.
- [19] Vincent, K. (2020). Radiometric assessment of natural radioactivity levels and radiation hazard indices for soil samples in Kericho County, Kenya.
- [20] Wanjala, E. M. (2016). Assessment of human exposure to natural source of radiation on the soil in Tongaren Constituency of Bungoma County, Kenya.
- [21] Kannan, V., Rajan, M. P., Iyengar, M. A. R., & Ramesh, R. (2002). Distribution of natural and anthropogenic radionuclides in soil and beach sand samples of Kalpakkam (India) using hyper pure germanium (HPGe) gamma ray spectrometry. *Applied Radiation and isotopes*, 57(1), 109-119.
- [22] Najam, L. A., Al-Dbag, S. T., Wais, T. Y., & Mansour, H. (2022). Radiogenic heat production from natural radionuclides in sediments of the Tigris river in Mosul City, Iraq. *International Journal of Nuclear Energy Science and Technology*, 15(3-4), 302-316.